



# GC Insights: Identifying conditions that sculpted bedforms - Human insights to build an effective AI

John K. Hillier<sup>1</sup>, Chris Unsworth<sup>2</sup>, Luke De Clerk<sup>3</sup>, Sergey Savel'ev<sup>3</sup>

<sup>1</sup>Geography and Environment, Loughborough University, Loughborough, LE1 3TU, UK

<sup>5</sup> <sup>2</sup>School of Ocean Sciences, Bangor University, Bangor, LL59 5AB, UK.

<sup>3</sup>Dept. Physics, Loughborough University, Loughborough, LE1 3TU, UK.

Correspondence to: John K Hillier (j.hillier@lboro.ac.uk)

Abstract. 42 survey participants demonstrate that it is visually possible to recognise the type of flow that created bedforms (e.g. sand dunes, riverbed ripples) from short distance-depth profiles, but this is much harder for individual forms. An interpreter's geoscience expertise does not help, indicating a machine learning or 'AI' algorithm might be trained well from the data alone, especially if multiple bedforms are used.

## **1** Introduction

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Environmental flows shape the surface they flow over. The variety of features produced (e.g. sand ripples on a beach), known as bedforms, reflect and preserve characteristics (e.g. speed, depth) of the flowing ice, water or air (Venditti, 2012; Bullard et

- 15 al., 2011; Storrar and Stokes, 2007). The relationships between bedform morphology and flow are contested where observation is extremely difficult, such as under ice-sheets (e.g. Clark et al., 2018; Hillier et al., 2018; Rose, 1987; King et al., 2009), and best understood for unidirectional water flow over sand in a laboratory setting, mimicking a river (e.g. Fig. 1a). Even in this idealised fluvial setting, it is difficult to construct a 1-to-1 link between bedform type (e.g. ripples or dunes) and specific flow conditions (Venditti, 2012; Froehlich, 2020). Illustratively, ripples have a higher aspect ratio (*H/L*, for height (*H*) and length
- 20 (L)) than dunes (e.g. Allen, 1968); yet the observational ranges overlap (Venditti, 2012; Yalin, 1972), creating uncertainty when attempting to link morphology with hydraulic conditions. Many variables related to hydraulics and/or the physics of sediment movement have been proposed to remove the overlap in bedform stability diagrams such as Fig. 1a (see Venditti, 2012). Only recently, has a distinct and non-overlapping zonation of bedform type and flow structure been developed using a quantity called shear velocity (Duran Vinet et al., 2019). Inverting this result may help realise the aspiration of developing a
- 25 means to reliably infer flow conditions from bedform morphology (e.g. Duran Vinet et al., 2019; Venditti, 2012; Myrow et al., 2018), which is often the only option (e.g. sedimentary structures preserving the geological past, Mars)(e.g. Ohata et al., 2017; Edgett and Lancaster, 1993).

Machine learning or 'AI' algorithms, such as Artificial Neural Networks (ANNs) offer an opportunity to examine this problem 30 as they do not assume simple (e.g. linear or 1-to-1) relationships between inputs and predicted variables (Wang et al., 2009;





Faruk, 2010). This has been attempted for experimental parameters (Froehlich, 2020), but not for bedform morphology. There may be unexploited quantitative morphological subtleties to categorise bedforms, or even to accurately position them on stability diagrams. This work examines the scope for using ANNs to distinguish the flow conditions in which relict bedforms originated by asking if the ability exists in non-artificial (i.e. human) intelligence for two particulars:

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Q1 - Is it possible to identify the environment (e.g. river, desert) of a bedform's genesis from its shape?

Q2 - In the fluvial environment, is it possible to distinguish flow conditions?

# 2 Method, Data & Ethics

An online survey (Supplementary Material) was conducted, initially at the 'Non-equilibrium flows and landforms' workshop (19th May 2021), with participation expanded using authors' close contacts (friends, colleagues, and family). For Q1, 40 participants attributed distance-depth profiles across 34 individual bedforms, and 13 bedform sequences ( $\geq$ 3 bedforms) to one of four environments (fluvial [river], glacial, marine, aeolian [desert]). For Q2, participants ranked three profiles according to flow strength (shear velocity), thrice for individual forms, and thrice for bedform sequences. Examples were provided to isolate visual shape analysis from prior knowledge (Fig. 1b), black and white profiles were used to exclude contextual clues (e.g. dataset characteristics, other features in the landscape), and the order of options (e.g. B, A, C) was shuffled for each participant 45 to prevent bias. Scale (e.g. metres) readily distinguishes environment without using bedform shape, so it was not given.

Aeolian data are ASTER (v2) across linear and transverse dune types from the Namib desert (Bullard et al., 2011), glacial are from near Lough Gara in Ireland (Hillier and Smith, 2008), fluvial are from four laboratory experiments (Expts. 1-4) of increasing shear velocity (Unsworth, 2015), and marine data are from the Irish Sea. Representative examples of individual 50 bedforms and sequences were manually selected from these datasets.

Ethical approval was given by the Ethics Review Sub-Committee at Loughborough University.

- ANN analysis to follow up the survey used a Multi-Layered Perceptron (MLP) with four hidden layers with 28, 56, 56 and 28 55 nodes, each with a ReLU activation function. Height (H) and width (W) fitted using the SWT algorithm (Hillier, 2008) and a frustum approximation (Hillier, 2006), see Fig. 1c & d, were input to predict the flow regime (experiment number). Groups consisted of five bedforms. Selection was random, and without replacement from a single time series until all bedforms in that time series were used. Weights and biases were updated using the Adam Optimiser of PyTorch using a loss function that 60 calculates the Mean Squared Error (MSE), all within a feedforward back-propagation approach.





# **3** Results

Of the 42 survey participants 25 self-identified as geoscientists, and 16 did not. For Q1, participants correctly identified the one of four environments (e.g. fluvial, aeolian) in which individual features originated 32% of the time, slightly if significantly (2-tailed t-test  $p \ll 0.01$ ) better than the 25% expected of guesswork. This rises to 51% for bedform sequences. For Q2

65 (2-tailed t-test,  $p \ll 0.01$ ) better than the 25% expected of guesswork. This rises to 51% for bedform sequences. For Q2, participants ranked entirely correctly 3 flow strengths (Expts. 1-3) for 46% of individual features, and 60% of sequences, much better than the 16% expected of guesswork (conservatively assuming no strength is repeated,  $p \ll 0.01$ ).

In none of the questions did geoscientists perform better than non-geoscientists, with mean percentages of correct answers being indistinguishable (2-tailed t-test, p > 0.05). The overall sentiment is encapsulated by one comment:

"I felt this was a geometrical exercise of recognising same patterns at different scales. I did not feel that my experience as an "expert" in bedforms really made any difference from, say, my son taking the test."

## **4** Discussion

- 75 The survey results clearly demonstrate it is possible to distinguish fluvial flow conditions from distance-depth data of the bed, and that an ANN should perform better if utilising sequences of bedforms rather than evaluating individuals in isolation. Geoscientists' *a priori* and contextual knowledge added little, indicating that training an ANN on these data alone should be productive.
- Morphologies from differing environments (e.g. glacial, fluvial) are often viewed as similar, indicators of analogous processes at work (e.g. Shaw, 1983), and modelled with identical equations (e.g. Fowler, 2002; Duran Vinet et al., 2019) or statistics (e.g. Hillier et al., 2016; Einstein, 1937). Several participants commented that their ability to distinguish environments might be to do with characteristics of the data (e.g. smoothness due to data resolution), not bedform shape, highlighting a potential pit-fall of training an ANN on raw distance-depth data. Another limitation is that ANNs performing pure pattern recognition need 1000s of training datasets (e.g. Bishop, 1996), which are not readily available in geoscience, and are apparently not needed by the survey's participants. Potentially, this requirement can be avoided by pre-processing to identify key aspects of

a geomorphological pattern (e.g. Shumack et al., 2020).

Shear velocity increases non-linearly across experiments 1 to 4. Fig. 1c-e present the outputs of a follow-up analysis into (i) the plausibility of building an ANN using only 4 time-series if pre-processing heights by fitting flat-topped cones to bedforms (Hillier, 2006, 2008), and (ii) the potential benefit of using short bedform sequences. Visually (Fig. 1d), the individual shapes (*H*, *W*) overlap between experiments, but the trends and averages over a number of points are distinctly different. This maps





directly to results of the ANN (Fig. 1e). Individual forms are weakly predicted (light grey,  $r^2 = 0.11$ ), but sub-sets of 5 bedforms more strongly so (grey,  $r^2 = 0.56$ ), particularly if very small bedforms present in all experiments (H < 0.5 cm) are excluded

95 (dark grey, r<sup>2</sup> = 0.80). Interestingly, simply taking the first four moments (mean, variance, skew, kurtosis) for arbitrarily located segments of the time-series, each 100 seconds long, yields better prediction (r<sup>2</sup> = 0.89). Thus, insights from the participants have contributed to building an effective AI to reliably infer flow conditions from bedform morphology, yet a trade-off may exist between the most discriminating ANN (e.g. from statistical properties - (e.g. Malinverno, 1988; Powell et al., 2016; Singh et al., 2011)) and ease of relating outputs to process understanding and traditional shape parameters (i.e. *H* and *W*). In future, in transitional, non-equilibrium conditions (e.g. see Myrow et al., 2018), ANNs may be key to disentangling forms and flow.

## **5** Audiences and wider application

Geoscience communication in this work was by engaging with non-geoscientists as an integral part of a research project. In doing so, it introduced geomorphology as a discipline and illustrated an active research question. The practical geoscientific

105 insight of wider interest is that derived properties of landscapes (e.g. shape parameters for several dunes) can assist training an AI where data are limited (i.e. observations of Earth, physical experiments).

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### References

Allen, J. R. L.: Their Relation to Patterns of Water and Sediment Motion, North Holland Publishing Company, Amsterdam, 433 pp., 1968.

Bishop, C. M.: Neural Networks for Pattern Recognition, Oxford University Press, 502 pp., 1996.

Bullard, J., White, K., and Livingstone, I.: Morphometric analysis of aeolian bedforms in the Namib Sand Sea using ASTER data, 36, 1534–1549, 2011.

Clark, C. D., Ely, J. C., Spagnolo, M., Hahn, U., Hughes, A., and Stokes, C. R.: Spatial orgainzation of drumlins, 43, 499– 513, https://doi.org/10.1002/esp.4192, 2018.

Duran Vinet, O. D., Adreiotti, B., Claudin, P., and Winter, C.: A unified model of ripples and dunes in water and planetary environments, 12, 345–350, https://doi.org/10.1038/s41561-019-0336-4, 2019.

Edgett, K. S. and Lancaster, N.: Volcaniclastic aeolian dunes: terrestrial examples and application to martian sands, 25, 271–297, https://doi.org/10.1006/jare.1993.1061, 1993.





125 Einstein, H. A.: Bedload tranport as a probability problem, in: Sedimentation, edited by: Shen, W. H., Colorado State University, Fort Collins, C1–C105, 1937.

Faruk, D. Ö.: A hybrid neural network and ARIMA model for water quality time series prediction, 23, 586–594, https://doi.org/10.1016/j.engappai.2009.09.015, 2010.

Fowler, A. C.: Evolution equations for dunes and drumlins, 96, 377-387, 2002.

130 Froehlich, D. C.: Neural Network Prediction of Alluvial Stream Bedforms, 146, 04020084, https://doi.org/10.1061/(ASCE)HY.1943-7900.0001831, 2020.

Hillier, J. K.: Pacific seamount volcanism in space and time, 168, 877-889, https://doi.org/10.1111/j.1365-246X.2006.03250.x, 2006.

Hillier, J. K.: Seamount detection and isolation with a modified wavelet transform, 20, https://doi.org/10.1111/j.1365-2117.2008.00382.x, 2008.

Hillier, J. K. and Smith, M.: Residual relief separation: digital elevation model enhancement for geomorphological mapping, 33, 2266–2276, https://doi.org/10.1002/esp, 2008.

Hillier, J. K., Kougioumtzoglou, I. A., Stokes, C. R., Smith, M. J., and Clark, C. D.: Exploring explanations of subglacial bedform sizes using statistical models, 11, e0159489, https://doi.org/10.1371/journal.pone.0159489, 2016.

140 Hillier, J. K., Benediktsson, Í. Ö., Dowling, T. P. F., and Schomacker, A.: Production and preservation of the smallest drumlins, https://doi.org/10.1080/11035897.2018.1457714, 2018.

King, E. C., Hindmarsh, R. C. A., and Stokes, C. R.: Formation of mega-scale glacial lineations observed beneath a west Antarctic ice stream, 2, 585–596, 2009.

Lewis, D. W. and McConchie, D. M.: Practical sedimentology, 2nd ed., Springer, US, 212 pp., 1994.

145 Malinverno, A.: Statistical Studies Of Seafloor Geomorphology, PhD Thesis, Lamont, 1988.

Myrow, P. M., Jerolmack, D. J., and Perron, J. T.: Bedform Disequilibrium, 88, 1096–1113, https://doi.org/10.2110/jsr.2018.55, 2018.

Ohata, K., Naruse, H., Yokokawa, M., and Viparelli, E.: New Bedform Phase Diagrams and Discriminant Functions for Formative Conditions of Bedforms in Open-Channel Flows, 122, 2139–2158, https://doi.org/10.1002/2017JF004290, 2017.

150 Powell, D. M., Ockelford, A., Rice, S. P., Hillier, J. K., Nguyen, T., Reid, I., Tate, N. J., and Ackerley, D.: Structural properties of mobile armors formed at different flow strengths in gravel-bed rivers, 121, https://doi.org/10.1002/2015JF003794, 2016.

Rose, J.: Drumlins as part of a glacier bedform continuum, in: Drumlin Symposium, edited by: Menzies, J. and Rose, J., Balkema, Rotterdam, 103–116, 1987.

Shaw, J.: Drumlin formation related to inverted melt-water erosional marks, 29, 461-479, 1983.

155 Shumack, S., Hesse, P., and Farebrother, W.: Deep learning for dune pattern mapping with the AW3D30 global surface model, 45, 2417–2431, https://doi.org/10.1002/esp.4888, 2020.





Singh, A., Lanzoni, S., and Wilcock, P. R.: Multiscale statistical characterization of migrating bed forms in gravel and sand rivers, 47, W12526, https://doi.org/10.1029/2010WR010122, 2011.

Southard, J. B. and Boguchwal, L. A.: Bed configurations in steady unidirectional water flow part 2. Synthesis of flume data, 60, 658–679, 1990.

Storrar, R. and Stokes, C. R.: A glacial geomorphological map of Victoria Island, Canadian Arctic., 3, 191–201, https://doi.org/10.1080/jom.2007.9710838, 2007.

Unsworth, C.: River Dunes in Unsteady Conditions, PhD, University of Hull, 2015.

Venditti, J. G.: Bedforms in Sand-Bedded Rivers, in: Treatise on Geomorphology, edited by: Shroder, J. and Wohl, E., 137– 165 162, 2012.

Wang, W., Chau, K., Cheng, C., and Qui, L.: A comparison of performance of several artificial intelligence methods for forecasting monthly discharge time series, 374, 294–306, https://doi.org/10.1016/j.jhydrol.2009.06.019, 2009.

Yalin, M. S.: Mechanics of Sediment Transport., Pergamon, Oxford., 1972.

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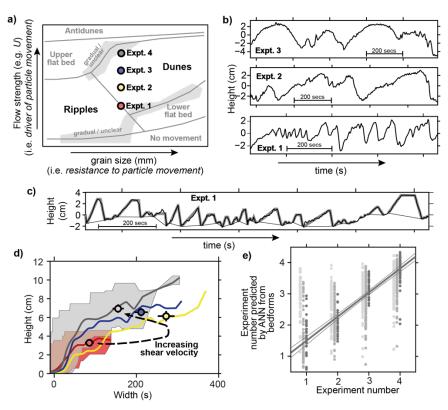


Figure 1: (a) Illustrative bedform stability 'phase diagram' for unidirectional fluvial (i.e. river) bedforms, synthesized from multiple sources (Ohata et al., 2017; Lewis and McConchie, 1994; Southard and Boguchwal, 1990). Main types considered here (i.e. ripples and dunes) are highlighted. Experiments 1-4 are positioned indicatively. (b) Time-series like those given unannotated to participants, i.e. one from each experiment 1-3, all scaled to the same dimensions. Horizontal axis is time because in the flume tank a stationary

i.e. one from each experiment 1-3, all scaled to the same dimensions. Horizontal axis is time because in the flume tank a stationary sensor recorded height as bedforms passed beneath it. (c) Example of how H and W are determined. Measured heights (thick black line) are processed using the SWT algorithm to identify bedforms, drawing a line beneath them (thin black line) then approximated as flat-topped cones (grey lines). SWT parameters as Hillier (2008). (d) Height-width relationships for the 4 experiments, with colours as in (a): lines are sliding means with W (Gaussian weights, width 60), shaded areas are full ranges for Expts. 1 & 4, and dots are the means  $(\pm 2\sigma)$  of upper quartile of the data when the small bedforms (i.e. H < 0.5 cm) are excluded. (e) Comparison of actual experiment number for out of sample prediction by the ANN using H and W: individual bedforms (light grey), subsets of 5 bedforms with (grey) and excluding (dark grey) small bedforms.